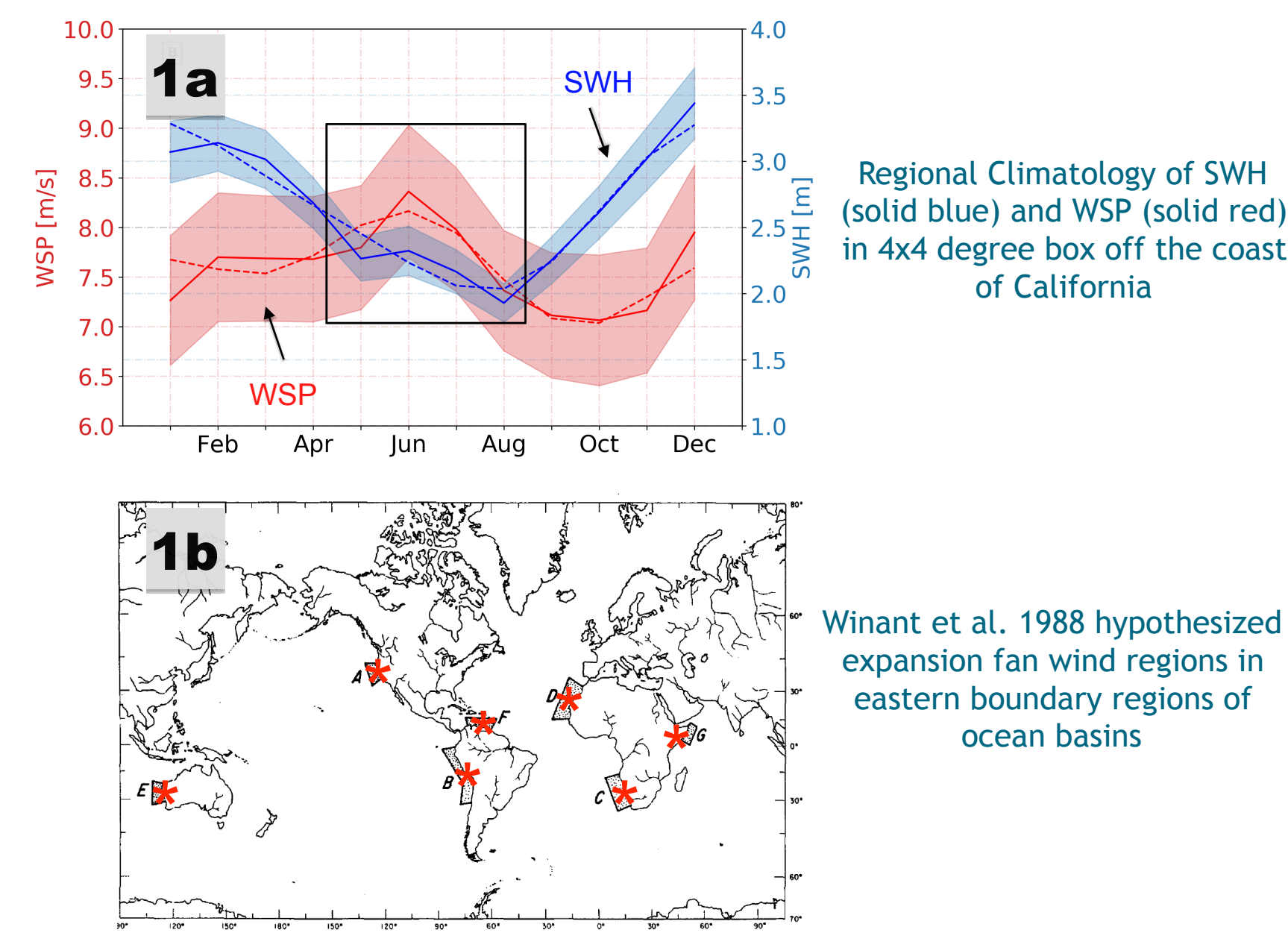


Motivation

Wind generated surface gravity waves are fundamental to our understanding of the interactions between the ocean and atmosphere. An appropriate metric for expressing quantitatively this interaction is **significant wave height (SWH)**. Remotely forced wind waves, generated by storm systems, dominate the wave field¹.

SWH of wind waves undergoes a sinusoidal annual cycle due to seasonal patterns in storm. However, local wind events can cause deviations from this annual cycle. Villas Bôas et al. analyzed a particular deviation from SWH's annual cycle off the California coast¹ due to the local wind event called **expansion fan winds² (EFWs)**. EFWs are identified by high wind speed (WSP) along shore winds and this deviation is characterized by an increase in SWH during the late spring to early summer months. This SWH anomaly has been linked to EFWs².



Similar local wind events has been hypothesized by Winant et al. to be present in other regions with similar coastal topography and atmospheric conditions as California (hereon referred as local wind anomaly regions (LWARs)). In the present work, we assess:

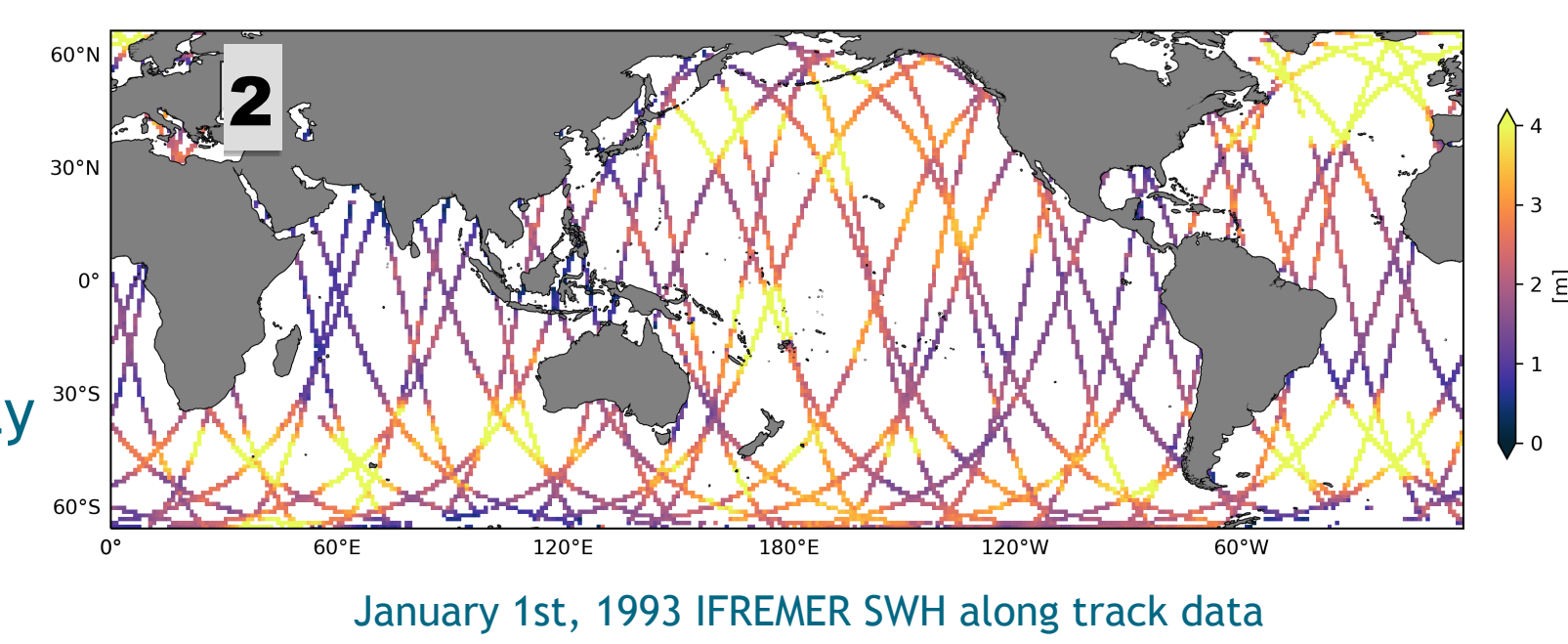
- Are deviations from the SWH annual cycle present in other LWARs?
- Is there a corresponding maximum in WSP climatology correlated with the SWH deviation?

Data and methods

Satellites measures SWH and WSP globally over large time periods. We used data from January 1st, 1993 to December 31st, 2015.

IFREMER SWH Satellite Altimetry Data

IFREMER's along track cross calibrated SWH altimetry data was collected from multiple near polar non-sun synchronous satellites binned to global 1x1 degree spatial and daily temporal resolution



Cross-Calibrated Multi-platform version 2 Wind Vector Product

CCMP2 is a gridded wind vector product consisting of a mixture of satellite remote sensed, in situ buoy, and modeled wind velocity data with global 0.25x0.25 degree spatial and 6 hourly temporal resolution

All analysis between IFREMER SWH and CCMP2 WSP is implemented with monthly averaged 1x1 degree spatial resolution data.

Annual and Semi-annual SWH and WSP Model

In order to analyze the seasonal cycle of SWH and WSP, we least-squares fitted SWH and WSP with annual and semi-annual sinusoidal cycles:

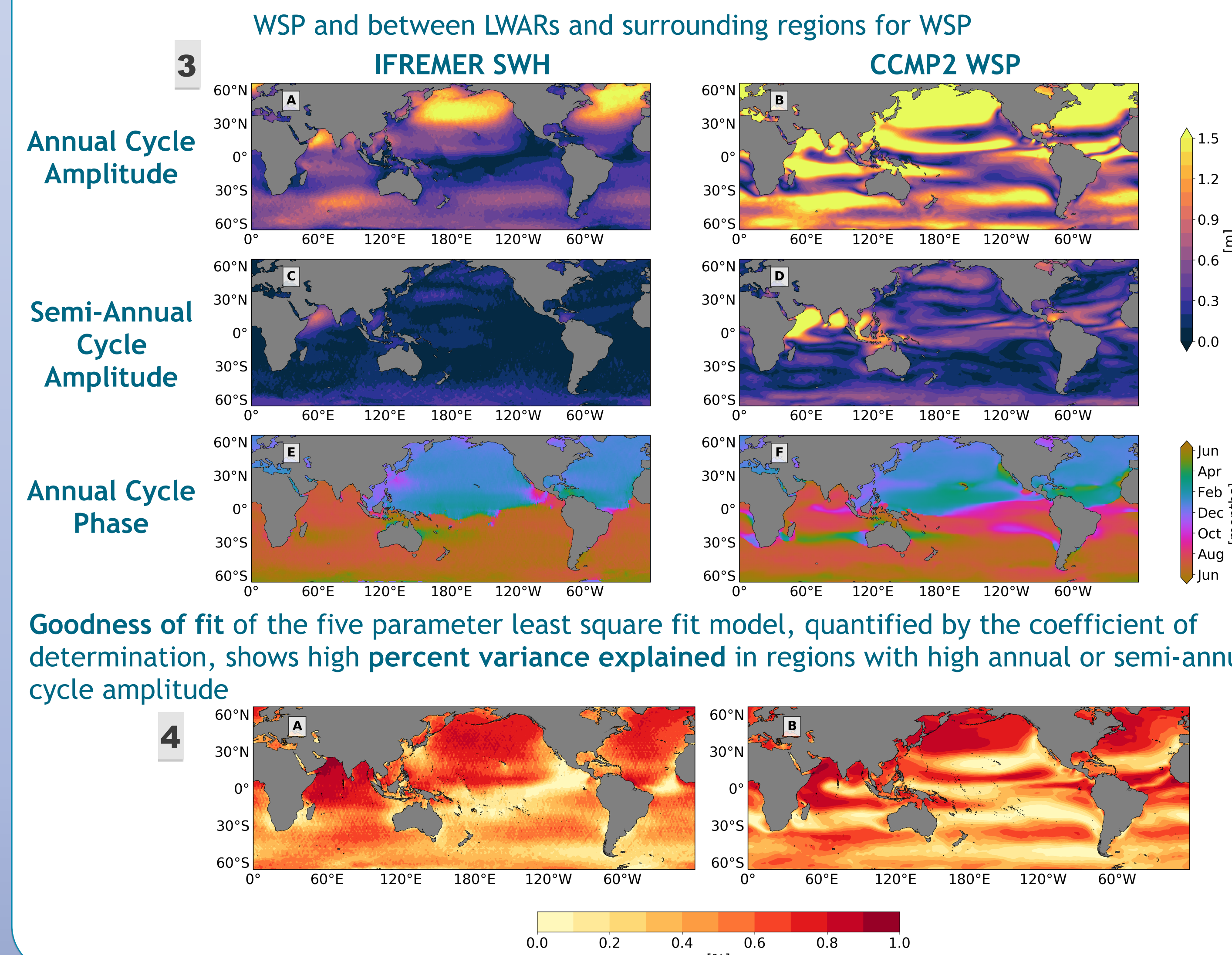
$$f(t) = a_0 + a_1 * \cos\left(\frac{2\pi t}{T}\right) + a_2 * \sin\left(\frac{2\pi t}{T}\right) + a_3 * \cos\left(\frac{4\pi t}{T}\right) + a_4 * \sin\left(\frac{4\pi t}{T}\right)$$

By fitting the regression curve along time series at each grid point across the globe for SWH and WSP data, the parameters of each annual cycle could be compared geographically.

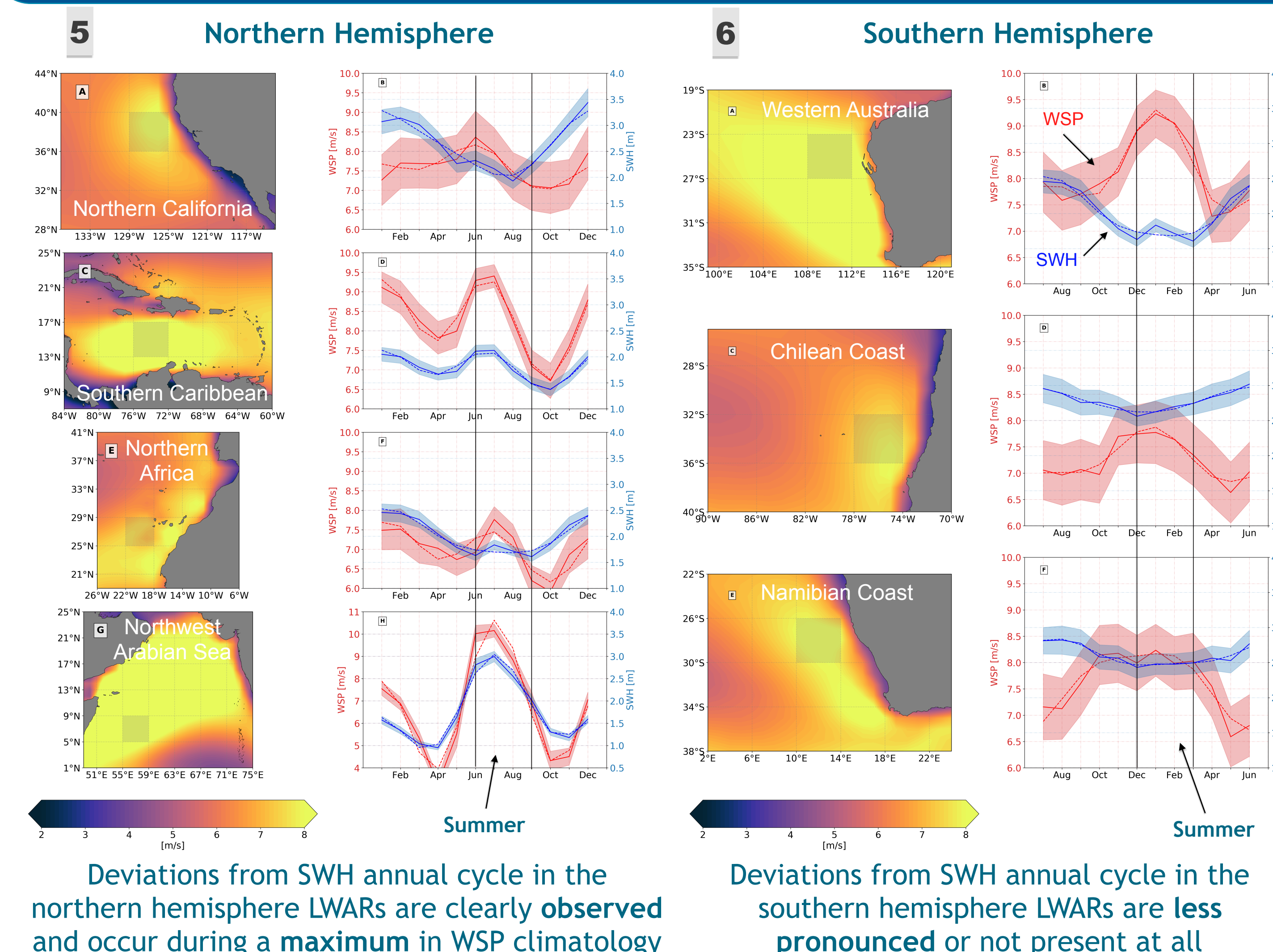
Parameters of Annual and Semi-annual Model

The parameters of the annual and semi-annual model reveal:

- High annual cycle amplitude in high latitudes of northern hemisphere
- High semi-annual cycle amplitude in monsoon and some LWAR regions
- Anti-correlation between northern and southern hemisphere for SWH and



Regional Climatology Analysis



The magnitude of the deviation from the SWH cycle is determined by the local conditions and characteristics of the wave field within the region

Probability of Swell: Consistency between Metrics

Wave Age

In order to justify that SWH during the spring months in these LWARs are locally forced, wave age can be used for separating growing seas from fully developed seas by quantifying the stage of development of the wave through the ratio criterion:

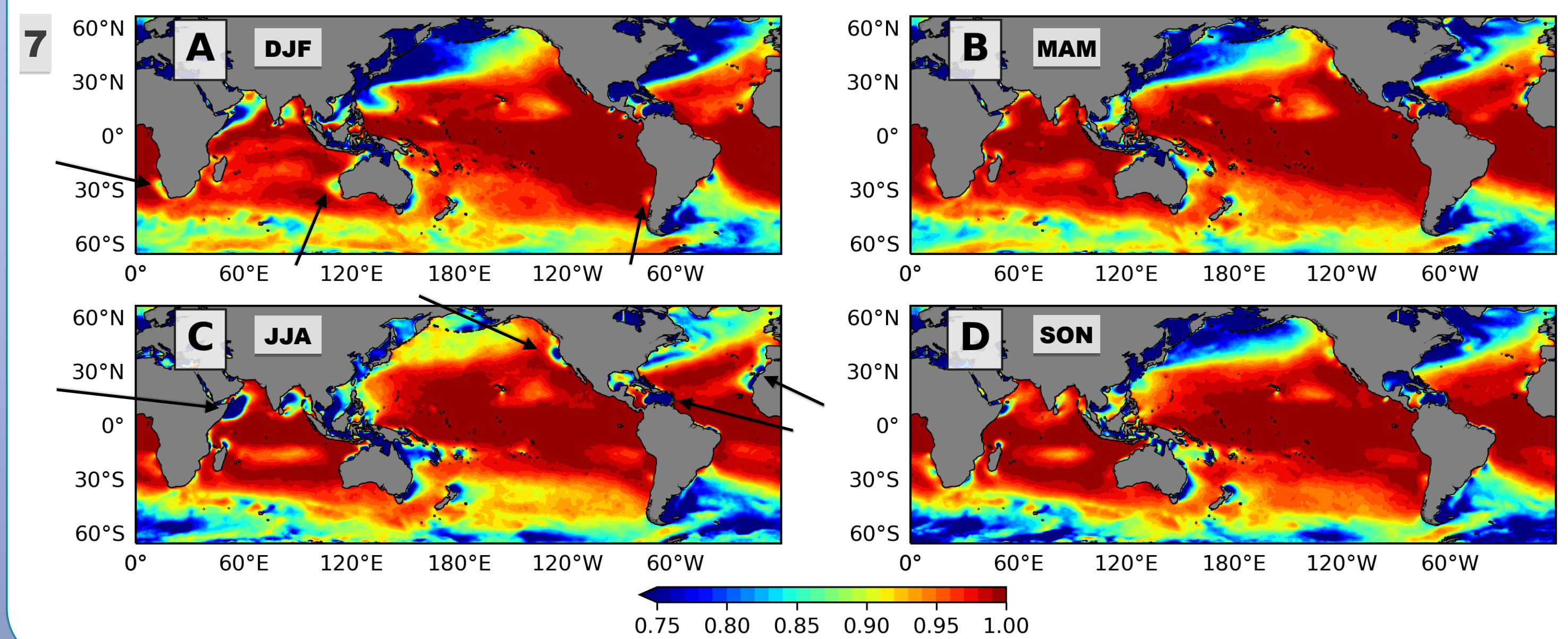
$$A = \frac{c_p}{U_{10}} \quad \text{where} \quad \begin{matrix} A > 1.2 \rightarrow \text{Remotely Forced Wave (Swell)} \\ A \leq 1.2 \rightarrow \text{Locally Forced Wave (Wind-seas)} \end{matrix}$$

Probability of Swell

Probability of swell quantifies the fraction of number of swell events (N_{swell}) to total number of wave events (N_{total}) over a time period using wave age as the separation criterion for wind-seas and swell events:

$$P_s = \frac{N_{swell}}{N_{total}} \quad \text{where} \quad N_{total} = N_{swell} + N_{wind}$$

In global seasonal progression maps of probability of swell below, LWARs contained waves that are categorized as locally forced during the spring and summer months for the respective hemisphere. In the southern hemisphere, this decrease in probability of swell is less than the northern hemisphere.



Final Remarks

- Deviations in the SWH seasonal cycle are present with high magnitude in the northern hemisphere and with low to near zero magnitude in the southern hemisphere due to close proximity to large southern ocean swells overwhelming locally forced waves
- Using wave age, probability of swell decreases significantly in all LWARs during the spring and summer months implying an increase in the amount of locally forced waves in LWAR. This is consistent with our hypothesis of local forced wave causing the deviation from the SWH annual cycle
- Future assessments will focus on using spectral data from WW3 to partition the wave field into swell and wind seas and plot regional climatologies of SWH for these partitions in order to further understand the dynamics between swell and wind-sea waves influencing the magnitude of the SWH deviation
- Understanding SWH climatology globally can help obtain realistic high-resolution ocean wave models.

References

- Villas Bôas, Ana B., et al. "Characterization of the deep water surface wave variability in the California Current region." *Journal of Geophysical Research: Oceans* 122.11 (2017): 8753-8769.
- Winant, C. D., et al. "The marine layer off northern California: An example of supercritical channel flow." *Journal of the Atmospheric Sciences* 45.23 (1988): 3588-3605.
- Stopa, Justin E., and Kwok Fai Cheung. "Intercomparison of wind and wave data from the ECMWF Reanalysis Interim and the NCEP Climate Forecast System Reanalysis." *Ocean Modelling* 75 (2014): 65-83.
- Hanson, Jeffrey L., et al. "Pacific hindcast performance of three numerical wave models." *Journal of Atmospheric and Oceanic Technology* 26.8 (2009): 1614-1633.
- Semedo, Alvaro, et al. "A global view on the wind sea and swell climate and variability from ERA-40." *Journal of Climate* 24.5 (2011): 1461-1479.